



The Laboratory of Atmospheric Chemistry





Aerosols



Definition: PM10 = Particles with aerodynamic diameter <10μm

Examples:

Diesel soot: ca. 0.1 µm



Ammonium sulfate: ca. 0.1 μ m



Pollen: 10 - 100 μm



Sea salt: 0.2 - 10 μm



Mineral dust: 0.2 - 10 µm







Size relationships



Diesel soot

100 Nanometer (nm) = 0.1 micrometer (μm) = 0.0001 millimeter (mm)



Pin head

2'000'000 nanometer (nm) = 2'000 micrometer (μm) = 2 Millimeter (mm)

Hot air balloon

40'000'000 micrometer (μm) = 40'000 millimeter (mm) = **40 Meter (m)**

In one cubiccentimeter of air: typically 10'000 particles





Aerosols

- Primary and secondary particles and size distributions
- Instrumentation
- Climate
- Health
- Source identification



Primary particle emissions in 2000 (Tg yr⁻¹)



	Northern Hemisphere	Southern Hemisphere	Global	Low	High	Source
Carbonaceous aerosols	• •	•	• •	•		
Organic Matter (0–2 μm	i)					
Biomass burning	28	26	54	45	80	Liousse <i>et al.</i> (1996), Scholes and Andreae (2000)
Fossil fuel	28	0.4	28	10	30	Cook et al. (1999),
						Penner et al. (1993)
Biogenic (>1µm)			56	0	90	Penner (1995)
Black Carbon (0–2 µm)						
Biomass burning	2.9	2.7	5.7	5	9	Liousse et al. (1996);
						Scholes and Andreae (2000)
Fossil fuel	6.5	0.1	6.6	6	8	Cooke <i>et al.</i> (1999);
Aircraft	0.005	0.0004	0.006			Penner <i>et al.</i> (1993)
Industrial Dust etc. (> 1 III	0.005 m)	0.0004	100	40	130	Wolf and Hidy (1997).
industrial Dust, etc. (> 1 µ)		100	40	150	Andreae (1995)
Sea Salt						Gong et al. (1998)
d< 1 µm	23	31	54	18	100	5
d=1-16µm	1,420	1,870	3,290	1,000	6,000	
Total	1,440	1,900	3,340	1,000	6,000	
Mineral (Soil) Dust ^b						
d< 1 µm	90	17	110	—		
$d=1-2\mu m$	240	50	290			
d=2-20µm	1,470	282	1,750	—		
Total	1,800	349	2,150	1,000	3,000	

^a Range reflects estimates reported in the literature. The actual range of uncertainty may encompass values larger and smaller than those reported here. ^b Source inventory prepared by P. Ginoux for the IPCC Model Intercomparison Workshop.







- Sulfate, homogeneous reaction: $SO_2 + \bullet OH + M \rightarrow HO^{\bullet}SO_2 + M$ $HO^{\bullet}SO_2 + O_2 + M \rightarrow HOO^{\bullet} + SO_3 + M$ $SO_3 + H_2O \rightarrow H_2SO_4$
- Sulfate, heterogeneous reaction: SO₂(g) ↔ SO₂(aq)
 SO₂(aq) + H₂O ↔ HSO₃⁻ (aq) + H₃O⁺(aq)
 Oxidation with H₂O₂, Ozon, NO₂, ...
- Nitrate, homogeneous reaction : $NO_2 + \bullet OH \rightarrow HNO_3$ $NH_3 (g) + HNO_3 (g) \leftrightarrow NH_4NO_3 (s)$
- Nitrate, heterogeneous reaction : $NO_2 + O_3 \rightarrow NO_3 (g) + O_2$ $NO_3 (g) + NO_2 + M \rightarrow N_2O_5 (g) + M$ $N_2O_5 (g) + H_2O (aq) \rightarrow 2 HNO_3 (aq)$
- Organics:

VOC + OH, O_3 , ... \rightarrow SOA (secondary organic aerosol)



Source strengths of sulfate and organic carbon (kg m⁻² hr⁻¹)





IPCC, 2001

Source strengths of black carbon, mineral dust and sea salt (kg m⁻² hr⁻¹)





0.0

0.2

0.1

60.0

1.0

0.5

120.0

2.0

5.0

180.0

10.0

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-120.0

0.01

0.02

-180.0

0.0

0.001

-60.0

0.05





IPCC, 2001

Size distributions of aerosol particles

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Figure 2.1. Schematic of aerosol sources, transformation mechanisms and sinks (modified after Whitby and Sverdrup (1980)).



Aerosol Size Distribution







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Traffic emissions: Influence of dilution temperature



Kittelson et al. (2000)



smallest particles are 10 times more abundant in the city compared to the country side and 100 times more abundant duirng the day than in the night



Bukowiecki et al., 2002







S. Pandis





Size distributions in Milan:

Evidence for primary and secondary particle formation







Secondary organic aerosol production







Terminology: Primary versus secondary

- Primary particles: directly emitted to the atmosphere
- Secondary particles: formed in the atmosphere by condensation (nucleation and growth) after chemical transformation
- How about oxidized primary particles ?
 aged primary (not secondary as Fuzzi et al. 2006 suggested)
- How about primary particles that evaporate on dilution and condense after oxidation (Robinson et al., 2007) ?
 secondary





Measurement Techniques

- Number
- Number size distribution
- Mass
- Optical properties
- Aerosol composition (off-line / on-line)







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Example: TSI CPC 3010 Lowest detectable diameter: D = 10 nm Maximum particle concentration: 10^4 cm⁻³

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Measuring the electrical mobility with a Differential Mobility Analyzers (DMA)



The center rod repels negative particles to the wall.

• Positively charged particles within a narrow mobility range

O Uncharged particles exit with the excess air.

A Para and

Measurable sizes: D=3-150 nm or 20-900 nm A scan is possible within 60-300 seconds

Principle:

- Defined electrical charging of the particles with radiocative source
- The aerosol flows laminarily through cylindric condensator
- An electrical field force the particles depending on their electrical mobility to move toward the inner electrode.
- Particles of a specific mobility $(b=v_p/E)$ are sucked through a gap at the inner electrode and are detected (typically by a CPC) afterwards.

$$b_{mech} = \frac{ln(r_{out}/r_{in})}{2\pi \cdot h} \cdot \frac{Q}{U \cdot ne} \quad b(D_p) = \frac{C_c(D_p)}{3\pi \eta \cdot D_p}$$

Q: total gas flow through DMA;

h, *r*_{in}, *r*_{out}: height, and inner and outer diameter of DMA cylinder;

U: applied voltage;

- e: elementary charge, 1.602e-19 As;
- n: number of elementary charges on aerosol particle.
- η : visosity of gas, air ~1.81e-5 Ns/m²;





Fast Analyzers:

EEPS Engine Exhaust Particle
Spectrometer. Scan time typ. 2 sec.
FPS Fast Particle Spectrometer
(similar to EEPS)
FMPS Fast Mobility Particle Sizer

(ambient air version of EEPS)

Engine Exhaust Particle Sizer Spectrometer (EEPS, TSI)

Fast Particle Spectrometer (DMS500, Cambustion)









Aerodynamic particle sizer







The PSI mobile laboratory











- Consistent picture : Nanoparticle concentrations <30 nanometers very high on highway
- In villages : much lower nanoparticle concentrations, in case of high wood burning contribution, higher volume concentration





Particulate mass

- Filters and gravimetric analysis
- On-line : e.g. Betameters, TEOM

Tapered element oscillating microbalance (TEOM)





"Lambert-Beer's Law" $I = I_0 e^{-b_{abs} \cdot x}$

Optical instruments e.g. light absorption by aethalomete

"Optical attenuation" ATN =
$$100 \times \ln\left(\frac{I_0}{I}\right) = 100 \times \ln\left(\frac{I_r}{I_s}\right)$$

For the time interval $\Delta t = t_2 - t_1$:

$$\Delta ATN = \ln\left(\frac{I_r(t_2)}{I_s(t_2)}\right) - \ln\left(\frac{I_r(t_1)}{I_s(t_1)}\right)$$

 I_0 S R

The attenuation coefficient (filtered aerosol)

The absorption coefficient (airborne aerosol)

$$b_{ATN} = \frac{A}{Q} \frac{\Delta ATN}{\Delta t}$$

Correction needed, e.g.

- Weingartner et al., 2003
- Arnott et al., 2005

 b_{abs}







The absorption exponent $\boldsymbol{\alpha}$





Absorption exponent α : a measure of the spectral variation in aerosol light absorption

- Example power law fit $(\lambda^{-2.0})$
- Wood burning^a $\lambda^{-1.8 \text{ to } -2.2}$
- **Traffic**, diesel soot^{a,b} $\lambda^{-1.0 \text{ to } -1.1}$

^a Kirchstetter et al. 2004

^b Schnaiter et al. 2003 & 2005

Enhanced UV-absorption due to the presence of wood smoke



Aerosol Chemical Analysis





Slide courtesy of Jose-Luis Jimenez





Examples of off-line analysis

- Ion chromatography (NH₄, NO₃, SO₄, organic acids)
- X-ray fluorescence, Particle Induced x-ray emission (PIXE), Inter Coupled Plasma mass spectrometry (ICP-MS), Neutron activation, Atomic Asorption Spectroscopy (AAS) (elemental analysis: K, S, Pb, Zn, ..)
- GC- / LC-MS (organic compounds: e.g. marker compounds hopanes, levoglucosan,)
- IR / UV / proton-NMR- spectroscopy : functional groups
- Mass spectrometry in general : isotope analysis, oligomers and more





On-Line Analysis

 Semi-online: EC-OC (separation of black/elemental carbon from organic carbon)







On-line analysis

- ATOF-MS
- Aerodyne Aerosol mass spectrometer







Example of Aerosol mass spectrometer measurements together with some black carbon measurements by an Aethalometer



Time resolution: minutes down to 6 seconds at low detection limits









Time Resolution Example -Fire Plumes from aircraft



3 Fires in the Yucatan Peninsula - MILAGRO field campaign







Direct and indirect aerosol effect on climate



Direct effect: Scattering and absorption of incoming sunlight by aerosol particles

Indirect effect:

The number concentration of cloud condensation nuclei (CCN) influences the cloud droplet size and thereby changes the cloud albedo and lifetime





Indirect aerosol effect



Large droplets → Weak reflection Small droplets → Strong reflection

Indirect effect

Number of CCN influences the droplet number and size (Twomey-Effect) and thereby the cloud albedo and lifetime.





'Ship tracks' visualise the indirect effect



Satellitenaufnahme (Wellenlänge: 3.7 µm)



Indirect aerosol effects





Figure 2.10. Schematic diagram showing the various radiative mechanisms associated with cloud effects that have been identified as significant in relation to aerosols (modified from Haywood and Boucher, 2000). The small black dots represent aerosol particles; the larger open circles cloud droplets. Straight lines represent the incident and reflected solar radiation, and wavy lines represent terrestrial radiation. The filled white circles indicate cloud droplet number concentration (CDNC). The unperturbed cloud contains larger cloud drops as only natural aerosols are available as cloud condensation nuclei, while the perturbed cloud contains a greater number of smaller cloud drops as both natural and anthropogenic aerosols are available as cloud condensation nuclei (CCN). The vertical grey dashes represent rainfall, and LWC refers to the liquid water content.





Climate forcing 2000 relative to 1750



Level of Scientific Understanding

IPCC (2001), www.ipcc.ch





Radiative forcing of climate between 1750 and 2005



No CO₂ time scale is given, as its removal from the atmosphere involves a range of processes that can span long time scales :







Aerosol direct radiative forcing – Comparison of different models





IPCC (2007)

Radiative forcing due to the Cloud albedo effect – Comparison of different models



IPCC (2007)





Total Aerosol Optical Depth (MODIS satellite) modelling





IPCC (2007)



Satellite retrieved AOD over Europe: Aerosols show high spatial variability







Lohmann, GRL, 2002

Evolution of particles in cloud: Bergeron-Findeisen process





Saturation Vapor Pressure (SVP) difference: SVP (**ice**) < SVP (**liquid**) \Rightarrow Flux of water vapor from liquid droplets to ice crystals

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The WBF mechanism converts many small supercooled drops to only few, large ice crystals, thus changing cloud radiative properties and enhancing precipitation.





Aerosols and Health

Diesel particles



- a health hazard

Aerosols is not a recent problem: The lung of 'Ötzi'



black: soot

red: soot

Source: www.ecocouncil.dk





Air pollution and mortality during London winter smog in 1952.



Adapted from Wilkins (1954)



Increased mortality correlates best with PM2.5 (fine particles)



Dockery et al. (1993)

PAUL SCHERRER INSTITUT **Epidemiology: clear relationship** LABOR FÜR TMOSPHÄRENbetween PM2.5 (or PM10) and mortality



CHEMIE





Different health effects of PM10

- Well established: Of all air pollutants, PM10 (and even more so PM2.5) show best correlation with increased mortality
- Increase of mortality with increase of PM10 by 10 μ g/m³ (Laden et al., 2000):
 - Traffic: 3.4%
 - Coal combustion: 1.1%
 - Mineral dust: ~0%
- ➔ Influence of
 - Chemical composition (metals, radicals, organic compounds, acidity)
 - Biological constituents (allergens, endotoxin)
 - Morphology (effect of asbestos)
 - Size distribution: smaller particles have greater surface area per unit mass
 - Number concentration
 - ...
- Mechanism not well known yet





Loss in life expectancy attributable to anthropogenic PM2.5 [months]



2000 2020 Current legislation

Loss in average statistical life expectancy due to identified anthropogenic PM2.5 Calculations for 1997 meteorology Provisional estimates with generic assumption on urban increment of PM

M. Amann, IIASA

2020

MTFR





Cytotoxicity of wood smoke from bad combustion: N. Klippel, Verenum





Verenum









Concentration of solid ammonium nitrate as a function of temperature





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Lanz et al., ACP (2007)





Plausibility of solution including 3 factors



Lanz et al., ES&T, accepted





Carbon apportionment using ¹⁴C analysis Estimation of fossil and non-fossil SOA contribution



Use of AMS analysis :

- wood burning 38%
- HOA 7%

Assumptions :

- only SOA, HOA and wood burning present

- OM/OC=2 for wood burning and SOA and OM/OC=1.2 for HOA

Lanz et al., accepted in ES&T





Smogchamber results of Carnegie Mellon



Figure 2 – Time series of aerosol mass (assuming $\rho = 1 \text{ g cc}^{-1}$) during an oxidation experiment and modeled primary mass assuming first order wall. Mass increases substantially when UV lights are turned on.

• Smog chamber results indicate that wood burning emissions is doubled after only 2 hours of chemistry



Figure 3 – Total aerosol mass (suspension + walls) during an experiment calculated with: 1) wall loss correction (with and without 'active' walls); 2) volume ratio. Also shown, against the right-hand axis, is the fractional contribution of the residual spectrum to the total AMS organic spectrum.



Roveredo in an Alpine valley, Januar 2005

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Average diurnal cycle of the carbonaceous material, OC/EC and wood burning versus traffic contributions



Sandradewi et al., ES&T, 2008



Aerosol mass spectra



Levoglucosan

Wood burner (emissions) chestnut, very inefficient burning

Night period in Roveredo in March, more than 80% of OC non-fossil

Average in Roveredo over the whole December

Mass spectra from a Motorway site in May



Alfarra et al. ES&T (2007)





SOA formation from TMB (at 50% RH)

Trimethylbenzene + NOx + light $\rightarrow \rightarrow$ Secondary Organic Aerosol







Kalberer et al., Science (2004)





Pattern of Zurich points to α -pinene rather than TMB (or other anthrop. precursors)

Baltensperger et al. Faraday Disc. (2005)

a.u.



Worldwide AMS measurements of the chemical composition





Zhang et al., GRL 2007





Models underestimate SOA



Volkamer et al. GRL 2006





SUMMARY

A significant progress was made in last years but many challenges remain:

- Better instrumentation (cheaper, smaller, more longterm, more specific, more precise, ...) is still needed
- Long-term chemical composition of aerosols is needed
- Secondary organic aerosol formation needs to be understood and implemented in models
- Health effects : more specific to size and chemical composition
- Climate : Indirect effects are not quantitatively understood